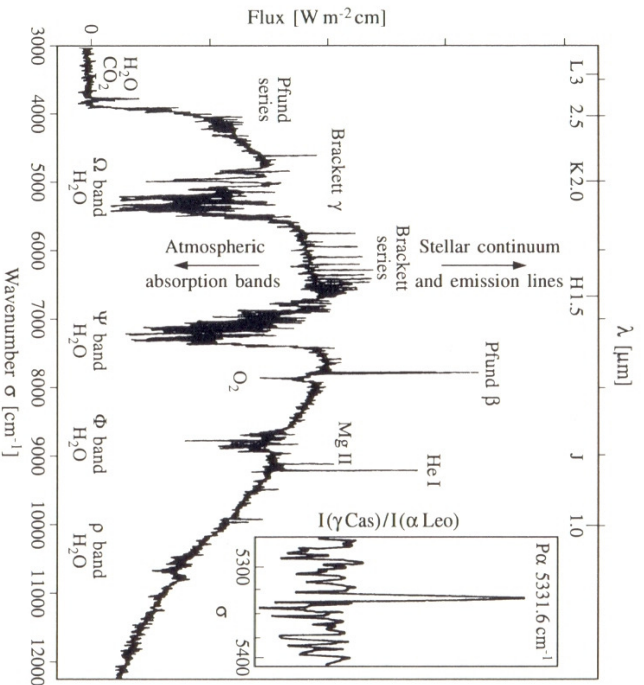


Astronomische Waarnemetechnieken (Astronomical Observing Techniques)

3rd Lecture: 22 September 2010



1. Atmospheric Layers
2. Absorption
3. Emission
4. Scattering, Refraction & Dispersion
5. Turbulence & Seeing

1. Atmospheric Layers

Assumption: atmosphere is in local radiative equilibrium and the composition is approximately constant.

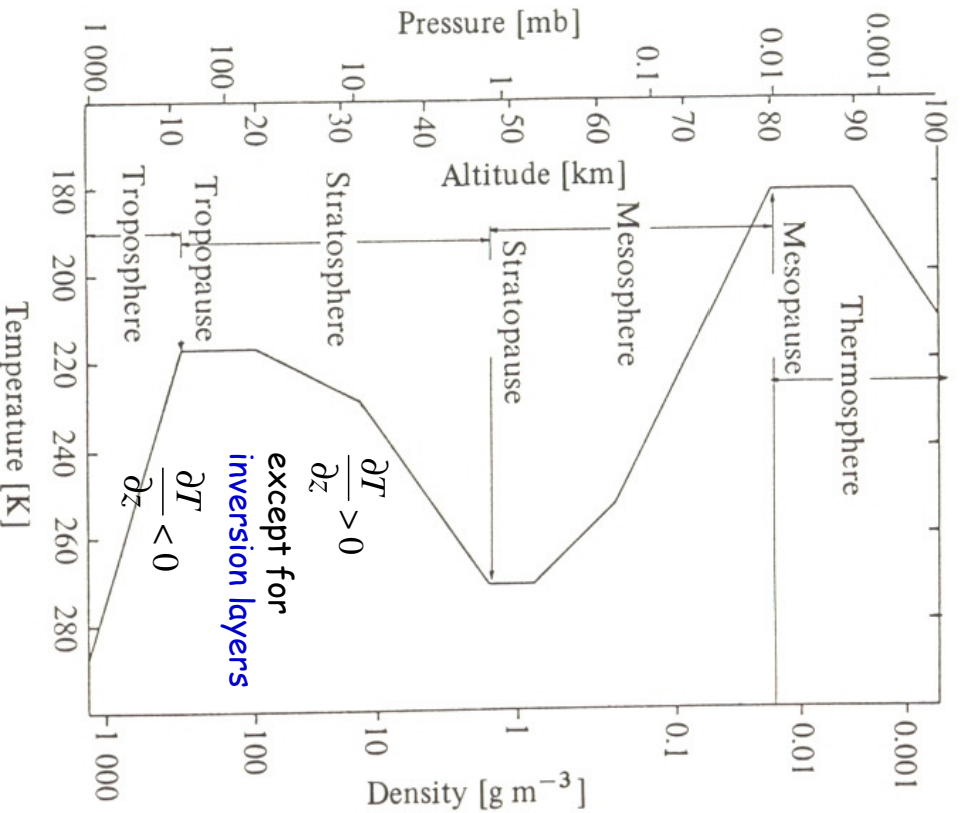
The structure can be described by three parameters:

- altitude z
- temperature $T(z)$
- density $\rho(z)$

The pressure $P(z)$ can be described by: $P(z) = P_0 e^{-\frac{z}{H}}$

where H = scale height ($H \sim 8\text{km}$ near ground).

Vertical Profile

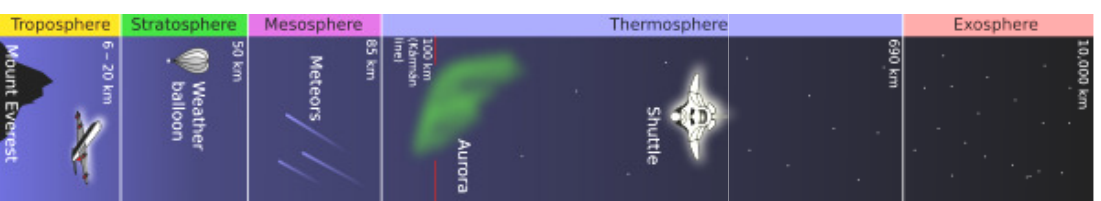


Constituents of the Atmosphere

- Main constituents: O_2 and N_2
relative constant proportions (78.1% N_2 , 20.9% O_2) up to 100 km
- **Water vapour** - causes very strong absorption bands
- **Ozone** - absorbs mainly in the UV
 - distribution depends on latitude and season
 - maximum concentration around 16 km height
- **CO_2** - important component for (mid)IR absorption
 - mixing independent of altitude (similar to N_2 , O_2)
- **Ions** - varies strongly with altitude and solar activity
 - relevant above 60km where reactions with UV photons occur:

$$O_2 + h\nu \rightarrow O_2^{*+} + e^- \quad \text{and} \quad O_2 + h\nu \rightarrow O^+ + O + e^-$$
 - electron showers along magnetic fields cause Aurora
 - at 100 - 300 km height: $n_e \sim 10^5 - 10^6 \text{ cm}^{-3}$

Ionosphere



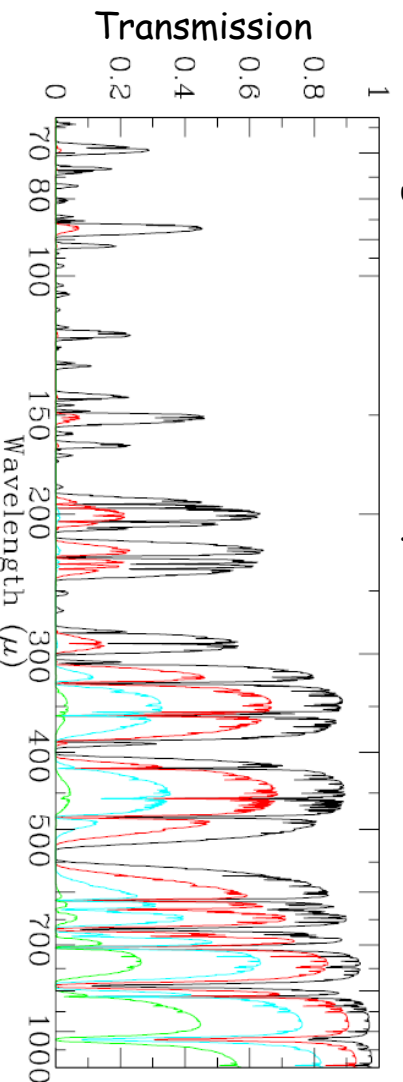
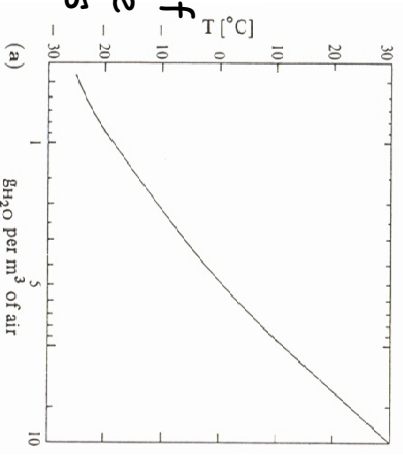
More on Water Vapor

The water vapor is a strong function of T and z .

The precipitable water vapor (PWV) is the depth of the amount of water in a column of the atmosphere if all the water in that column were precipitated as rain.

The amount of PWV above an altitude z_0 is:

$$w(z_0) = \int_{z_0}^{\infty} N_{H_2O} dz, \quad \text{where} \quad N_{H_2O} [\text{m}^{-3}] = 4.3 \times 10^{25} \frac{P}{P_0} \frac{T}{T_0} r(z)$$



Scale height for PWV is only ~ 3 km \rightarrow observatories on high altitudes

2. Absorption of Radiation

Atomic and molecular transitions that cause absorption features:

- pure rotational molecular transitions: H_2O , CO_2 , O_3 ,
- rotation-vibrational molecular transitions: CO_2 , NO , CO
- electronic molecular transitions: CH_4 , CO , H_2O , O_2 , O_3 , OH
- electronic atomic transitions: O , N , ...

The attenuation at altitude z_0 is given by:

$$\frac{I(z_0)}{I_0(\infty)} = \exp \left[-\frac{1}{\cos \theta} \sum_i \tau_i(\lambda, z_0) \right]$$

for i absorbing species with an optical depth of $\tau_i(\lambda, z_0) = \int_{z_0}^{\infty} r_i(z) \rho_0(z) \kappa_i(\lambda) dz$

(θ is the zenith distance; κ is the absorption coefficient; ρ_0 is the mass density of air).

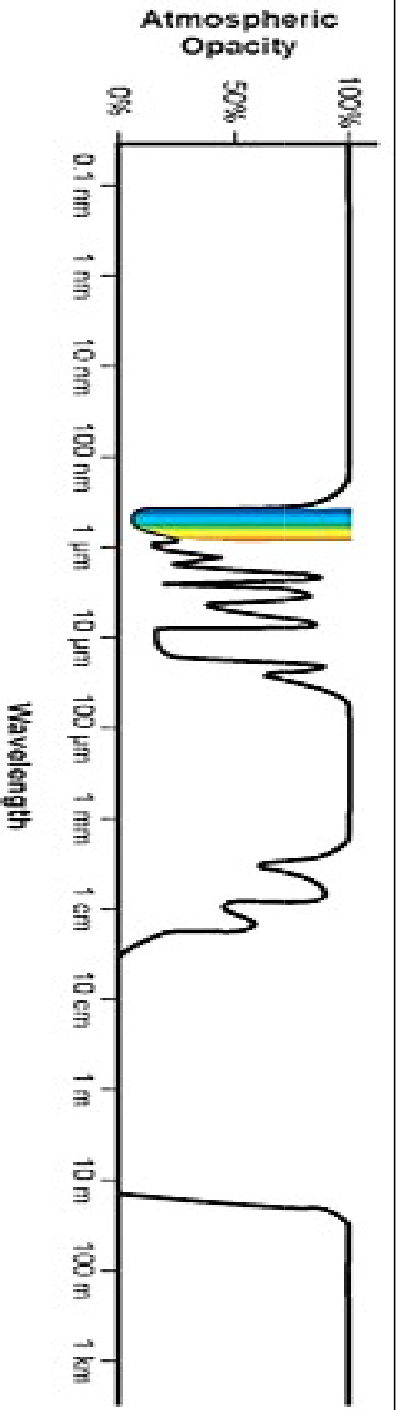
Atmospheric Bands

Two cases of absorption:

- total** absorption → atmospheric **transmission windows**
- partial** absorption → reduced transmission due to narrow **telluric*** absorption features

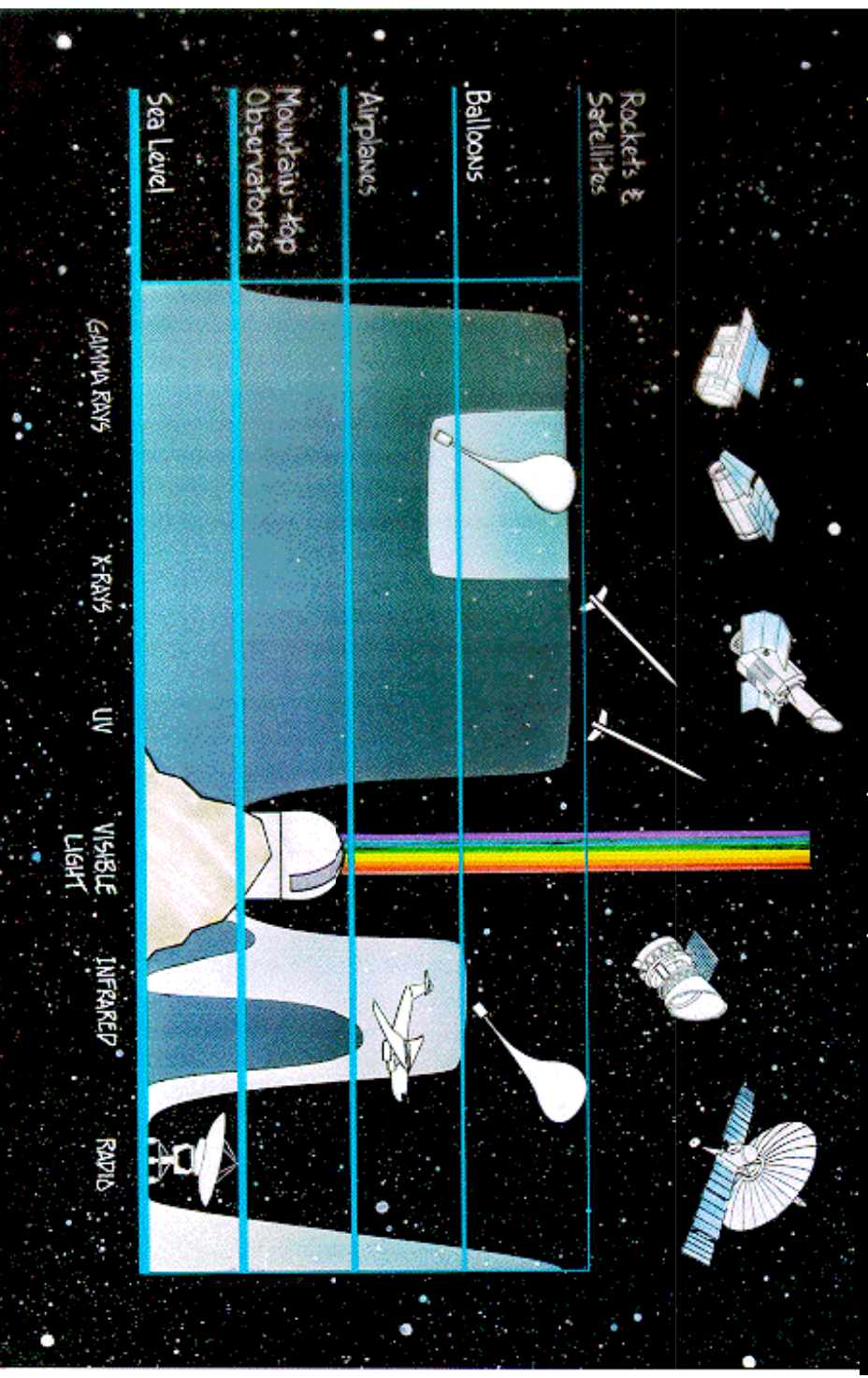
**Telluric = related to the Earth; of terrestrial origin*

The atmospheric opacity defines the **atmospheric transmission bands** - and thus the wavelengths that are accessible to observations



Ground based astronomy is limited to visible, near/mid-IR and radio wavelengths.

Space astronomy provides access to γ-rays, X-rays, UV, FIR, sub-mm



Side note: SOFIA

FIR/sub-mm astronomy is also possible from airplanes, e.g. the **Stratospheric Observatory for Infrared Astronomy (SOFIA)**

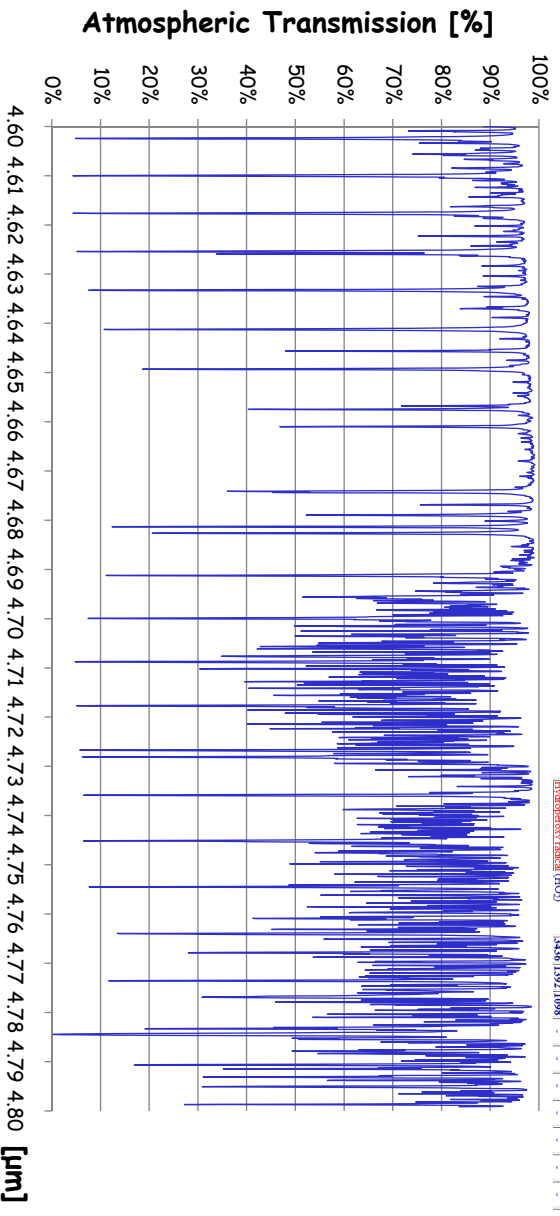


Side note: HITRAN

The HITRAN'2004 Database contains 1,734,469 spectral lines for 37 different molecules.

<http://cfa-www.harvard.edu/hitran/>

Molecule	13	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31	32
Acetylene (C ₂ H ₂)	3374	1974	3289	612	729	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Ammonia (NH ₃)	3337	950	3444	1627	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Carbon dioxide (CO ₂)	1388	667	2349	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Carbon monoxide (CO)	2143	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Carbon tetrachloride (CCl ₄)	464	217	799	316	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Carbon fluoride (CF ₄)	1945	965	582	1243	619	714	-	-	-	-	-	-	-	-	-	-	-	-	-
Carbonyl sulfide (OCS)	859	520	2062	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Chlorine oxide (ClONO ₂)	1737	1263	809	780	563	435	262	711	120	-	-	-	-	-	-	-	-	-	-
Chlorine oxide (ClO)	884	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Ethylene (C ₂ H ₄)	2954	1388	895	289	2886	1379	2069	1468	823	2085	1469	822	-	-	-	-	-	-	-
Formaldehyde (H ₂ CO)	3006	1623	1543	1023	3103	3238	949	943	3106	526	2989	1444	-	-	-	-	-	-	-
Formic acid (HCOOH)	2182	1746	1500	2843	1249	1167	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen fluoride (HF)	3570	3943	1770	1387	1220	1106	625	1003	638	-	-	-	-	-	-	-	-	-	-
Hydrogen chloride (HCl)	2559	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen cyanide (HCN)	3311	713	2097	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen cyanide (HNC)	3903	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen iodide (HI)	2230	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen peroxide (H ₂ O ₂) (t=1)	3593	1396	866	259	3560	1246	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydrogen peroxide (H ₂ O ₂)	2615	1183	2526	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-
Hydroxyl radical (OH)	3436	1392	1098	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-	-



3. Atmospheric Emission

A. Fluorescent Emission

Fluorescence = recombination of electrons with ions.

The recombination probability is low; takes several hours → night time

- Produces both continuum + line emission = airglow
- Occurs mainly at ~ 100 km height
- Main sources of emission are: O I, Na I, O₂, OH (←NIR), H

The emission intensity is measured in Rayleigh:

$$1 \text{ Rayleigh} = 10^6 \text{ photons cm}^{-2} \text{ s}^{-1} \text{ sr}^{-1} = \frac{1.58 \cdot 10^{-11}}{\lambda[\text{nm}]} \text{ W cm}^{-2} \text{ sr}^{-1}$$

B. Thermal Emission

Up to 60 km is the atmosphere in local thermodynamic equilibrium (LTE), i.e., the excitation levels are thermally populated.

Calculating the specific energy received requires a full radiative transfer calculation (see below), but for $\tau \ll 1$ one can use the approximation:

$$I_{\lambda}(z) = \tau_{\lambda} B_{\lambda}(\bar{T}) \frac{1}{\cos \theta}$$

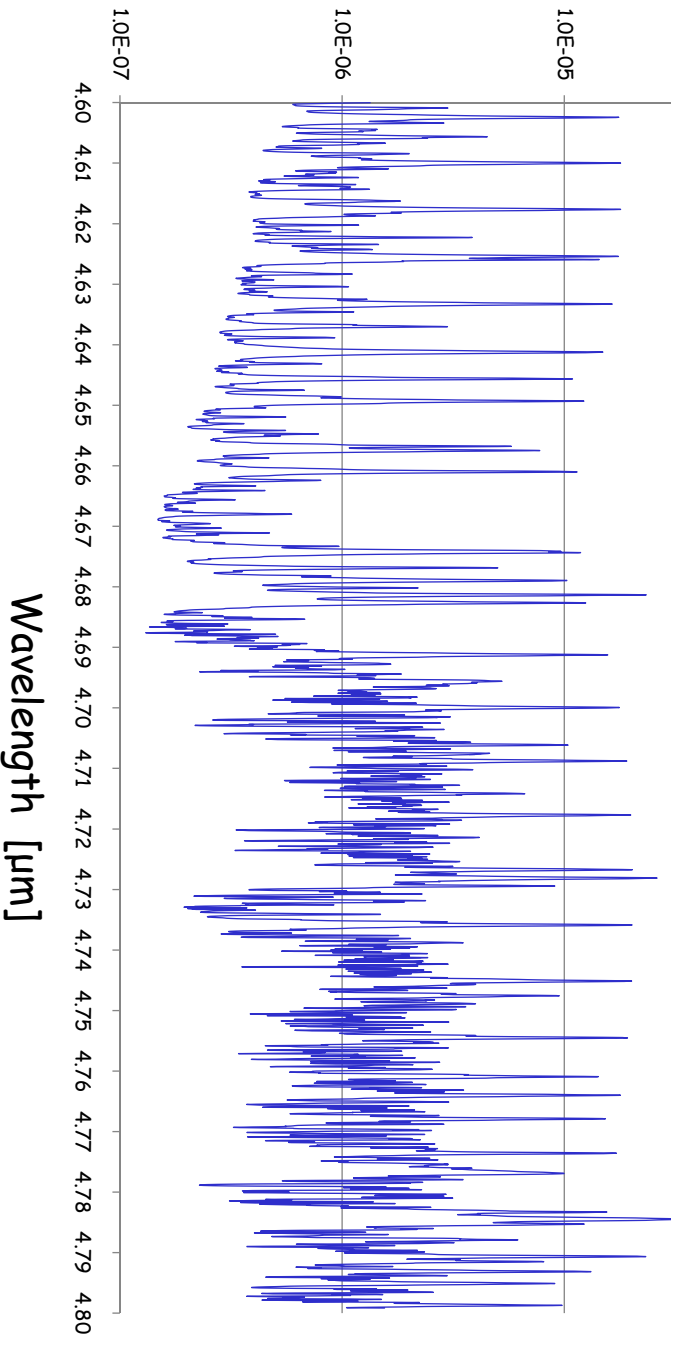
where $B(\bar{T})$ is the Planck function at the mean temperature of the atmosphere.

For $\bar{T} = 250 \text{ K}$ and $\theta = 0$:

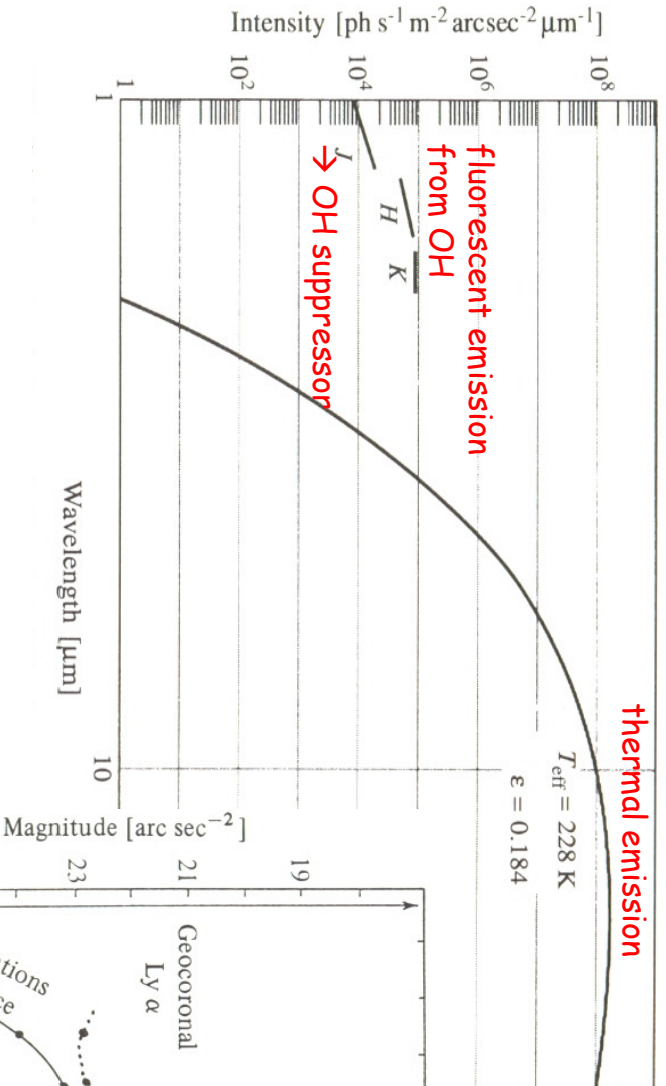
Spectral band (cf. Sect. 3.3)	L	M	N	Q
Mean wavelength [μm]	3.4	5.0	10.2	21.0
Mean optical depth τ	0.15	0.3	0.08	0.3
Magnitude [arcsec ⁻²]	8.1	2.0	-2.1	-5.8
Intensity [Jy arcsec ⁻²] ^a	0.16	22.5	250	2100

Exact Solution: full Radiative Transfer

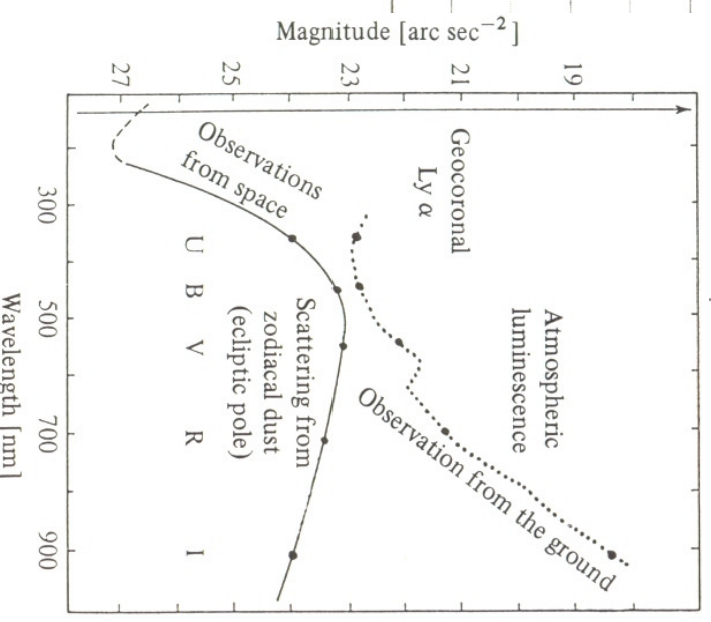
Atmospheric Emission [$\text{W}/\text{cm}^2/\mu\text{m}/\text{sr}$] (HITRAN)



Fluorescent and Thermal Emission



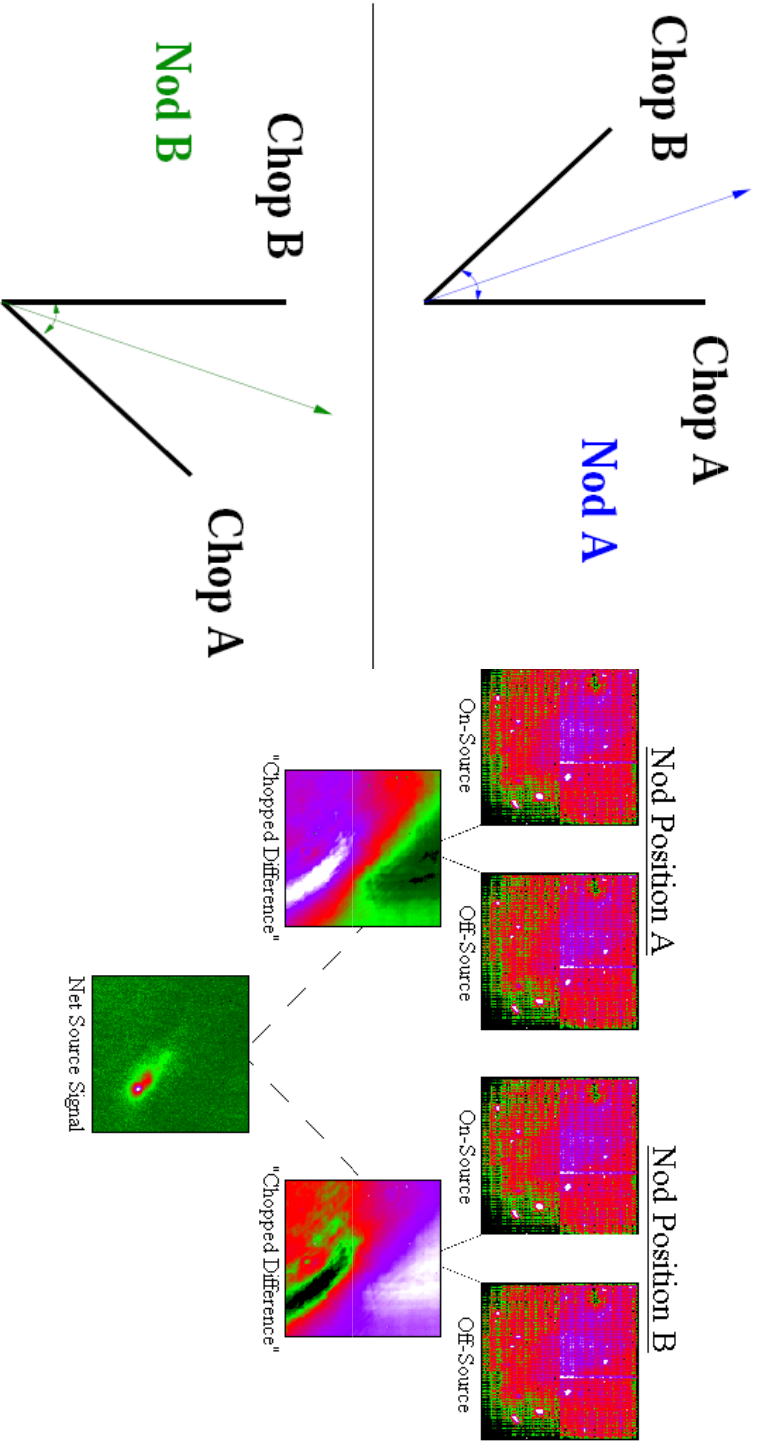
Sky brightness per arcsecond.
Important as even an unresolved point source has a finite angular diameter when viewed through a telescope.



Chopping & Nodding

High background: Poissonian photon shot noise + spatially & temporally varying fluxes + instrumental drifts/artefacts → chopping/nodding

Object *



4. Scattering, Refraction and Dispersion

A. Scattering by Air Molecules

Molecular scattering in the visible and NIR is Rayleigh scattering given by:

$$\sigma_R(\lambda) = \frac{8\pi^3}{3} \frac{(n^2 - 1)^2}{N^2 \lambda^4}$$

where N is the number of molecules per unit volume and n is the refractive index of air ($n-1 \sim 8 \cdot 10^{-5} P/T$).

Remember, Rayleigh scattering is **not isotropic**: $I_{scattered} = I_0 \frac{3}{16\pi} \sigma_R (1 + \cos^2 \theta) d\omega$

B. Aerosol Scattering

Aerosols (like sea salts, hydrocarbons, volcanic dust) are much bigger than air molecules → Rayleigh scattering does *not* apply.

Instead, scattering is described by **Mie's theory** (from classical electrodynamics, using a "scattering efficiency factor" Q):

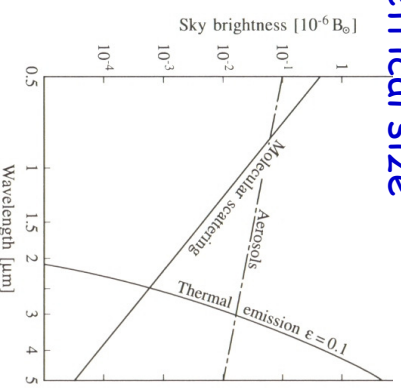
$$Q_{\text{scattering}} = \frac{\sigma_M}{\pi a^2} = \frac{\text{scattering cross section}}{\text{geometrical cross section}}$$

If $a \gg \lambda$ then $Q_{\text{scattering}} \sim Q_{\text{absorption}}$ and:

- the scattered power is equal to the absorbed power
- the effective cross section is twice the geometrical size

If $a \sim \lambda$ then $Q_s \propto 1/\lambda$ (for dielectric spheres):

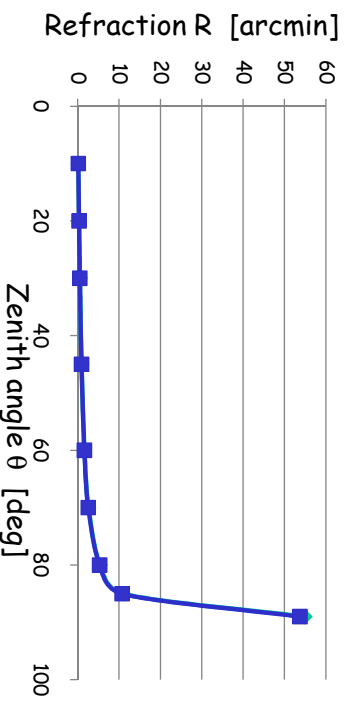
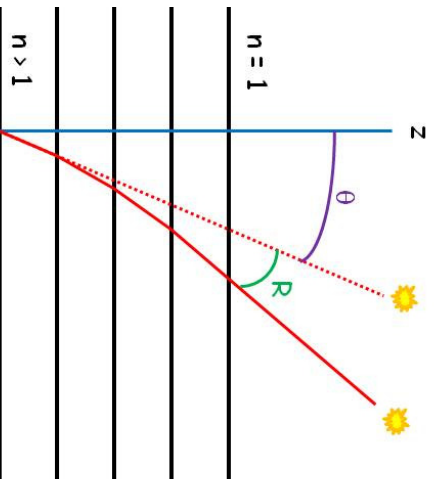
- the scattered intensity goes with $1/\lambda$



Atmospheric Refraction

Due to **atmospheric refraction**, the *apparent* location of a source is significantly altered (up to half a degree near the horizon) → telescope pointing model.

$$\text{Refraction } R = (n(\lambda) - 1) \tan \theta$$



Note that the refractive index of air depends on the wavelength λ :

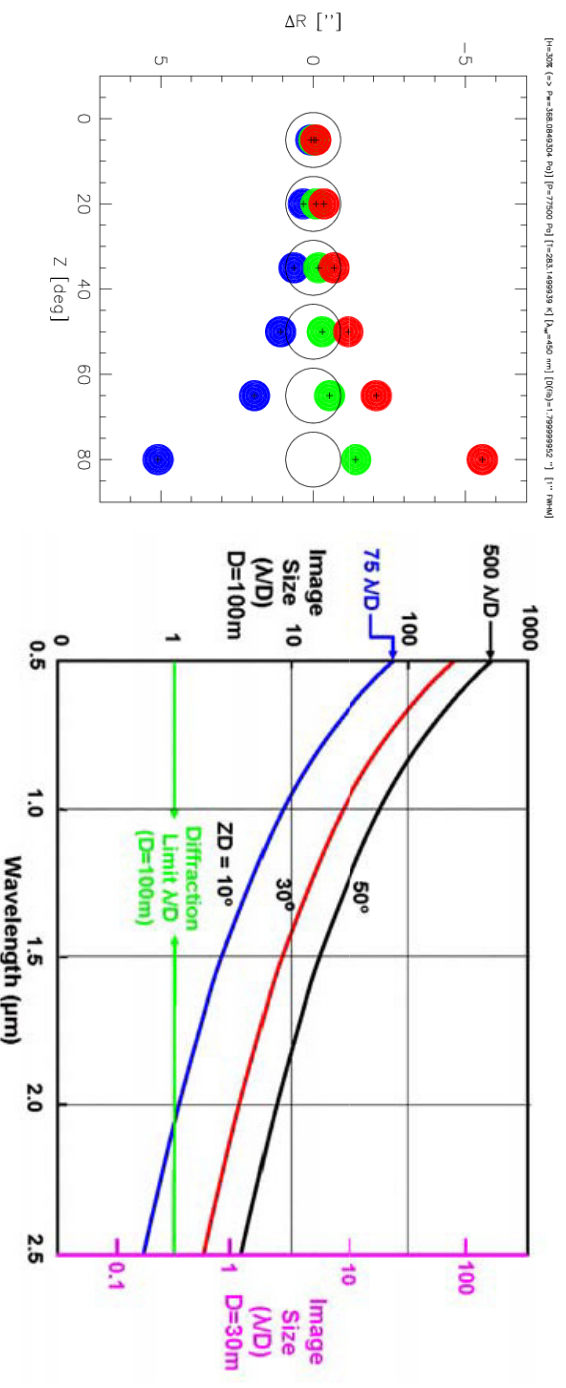
$$[n(\lambda) - 1] \times 10^6 = 64.328 + \frac{29498.1}{\lambda_0^2} + \frac{255.4}{\lambda_0^4} + \frac{146}{\lambda_0^6} - \frac{1}{\lambda_0^2} + \frac{41}{\lambda_0^4} - \frac{1}{\lambda_0^6}$$

(valid for dry air, 1 atm pressure, $T \sim 290\text{K}$ and λ_0 in [μm]).

Atmospheric Dispersion

Dispersion: The elongation of points in broadband filters due to $n(\lambda)$ [→ “rainbow”].
 The magnitude of the dispersion is a strong function of airmass and wavelength.

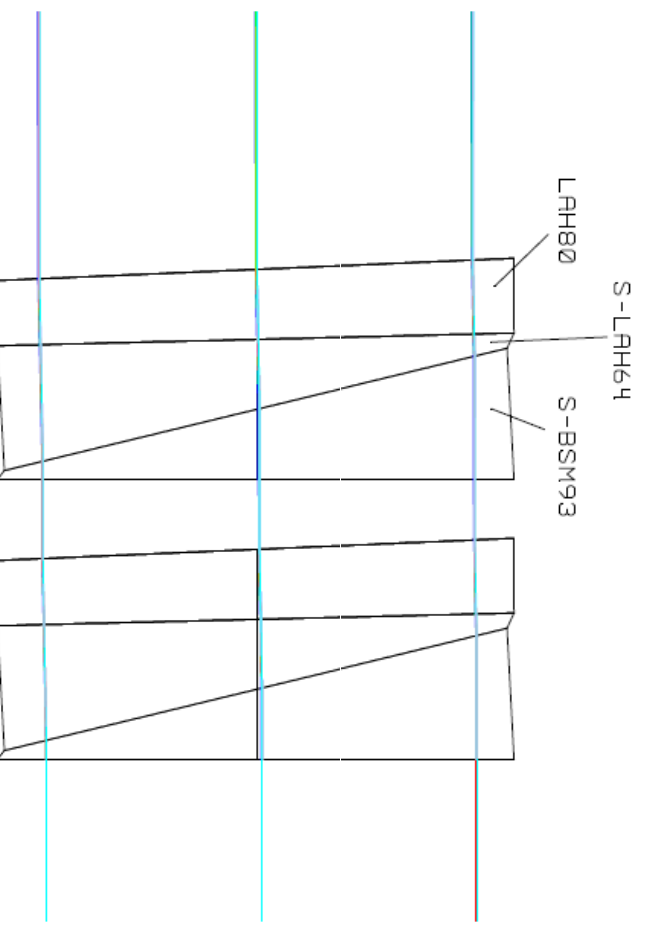
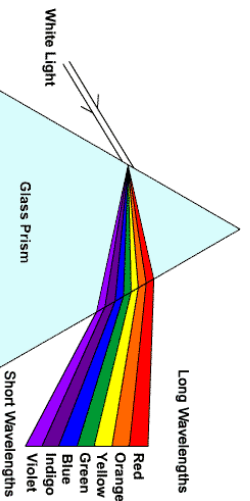
No problem is dispersion $\propto \lambda/D \leftarrow$ o.k. for small or seeing limited telescopes, but big problem for large diffraction limited telescopes



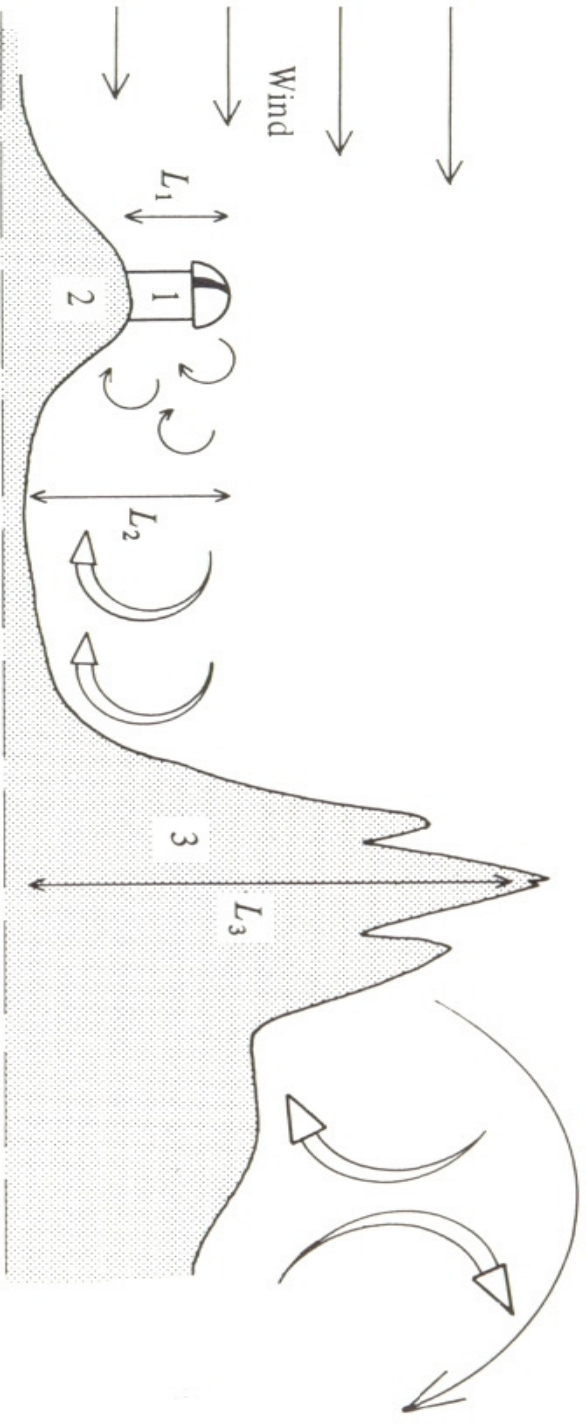
Atmospheric Dispersion Corrector

To counterbalance atmospheric dispersion use:

1. a refractive element (e.g., prism)
2. a second prism (different material with different dispersion) to maintain the optical axis
3. use a second (identical) double prism assembly to adjust the strength of the correction for different zenith angles.



5. Atmospheric Turbulence



The scales L_1 , L_2 , L_3 are characteristic of the outer (external) scales of turbulence caused by the wind around the obstacles 1, 2, 3.

The Reynolds Number

Turbulence develops in a fluid when the Reynolds number Re

$$Re = \frac{\rho VL}{\mu} = \frac{VL}{\nu}$$

exceeds a critical value.

V is the flow velocity

μ is the dynamic viscosity

ν the kinematic viscosity of the fluid ($\nu_{\text{air}} = 1.5 \cdot 10^{-5} \text{ m}^2 \cdot \text{s}^{-1}$)

L the characteristic length, e.g. a pipe diameter.

At $Re \sim 2200$ the transition from **laminar** to **turbulent** flow happens.

Example: wind speed $\sim 1 \text{ m/s}$, $L = 15 \text{ m} \rightarrow Re = 10^6 \rightarrow$ turbulent!

The Power Spectrum of Turbulence

The kinetic energy of large scale ($\sim L$) movements is gradually transferred to smaller and smaller scales, down to a minimum scale length l_0 , at which the energy is dissipated by viscous friction.

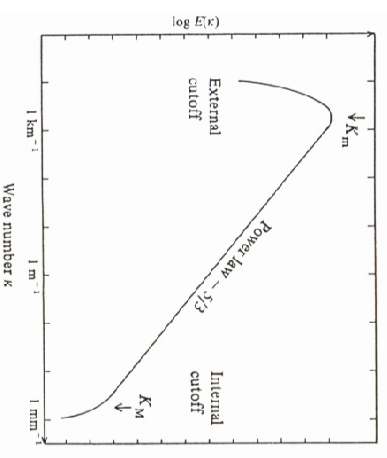
The local velocity field can be decomposed into spatial harmonics of the wave vector κ .

The reciprocal value $1/\kappa$ represents the scale under consideration.

The mean 1D spectrum of the kinetic energy, or Kolmogorov spectrum, is:

$$E(\kappa) \propto \kappa^{-5/3}$$

where l_0 is the inner scale, L_0 the outer scale of the turbulence, and $L_0^{-1} < \kappa < l_0^{-1}$



Air Refractive Index Fluctuations

Winds mix layers of different temperature \rightarrow fluctuations of temperature $T \rightarrow$ fluctuations of density $\rho \rightarrow$ fluctuations of refractive index n .

Of interest: difference between $n(r)$ at point r and $n(r+\rho)$ at a nearby point. The variance of the two values is given by:

$$D_n(\rho) = \langle [n(r) - n(r+\rho)]^2 \rangle = C_n^2 \rho^{2/3}$$

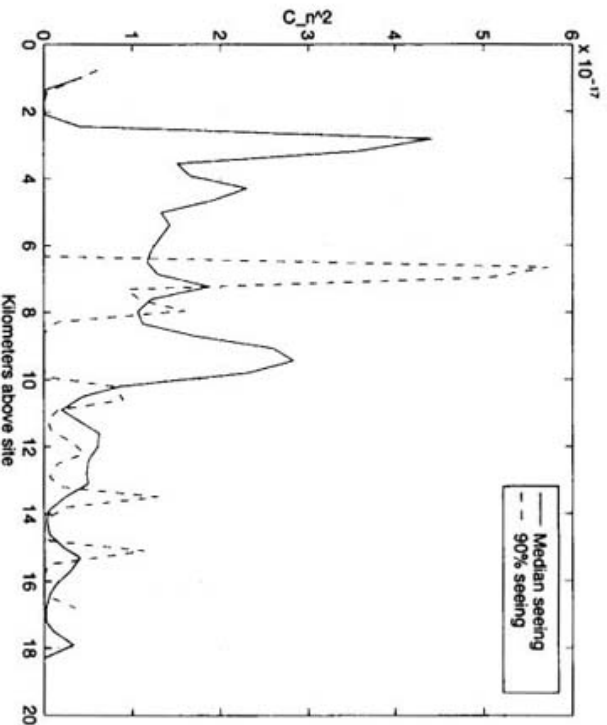
where $D_n(\rho)$ is the index structure function and C_n^2 is the index structure coefficient or structure constant of the refractive index.

Air Refractive Index Fluctuations (2)

Usually, one is only interested in the *integral* of fluctuations along the line of sight: $C_n^2 \cdot \Delta h$.

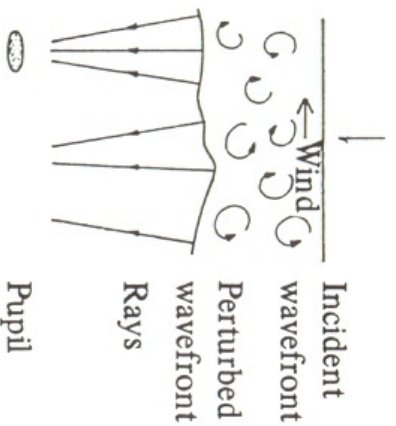
Typical value: $C_n^2 \cdot \Delta h \sim 4 \cdot 10^{-13} \text{ cm}^{1/3}$ for a 3 km altitude layer

But: there are always several layers of turbulence



Median seeing conditions on Mauna Kea are taken to be $r_o \sim 0.23$ meters at 0.55 microns. The 10% best seeing conditions are taken to be $r_o \sim 0.40$ meters. Figure taken from a paper by Ellerbroek and Tyler (1997).

Image Degradation by the Atmosphere



Scintillation - the energy received by the pupil varies in time

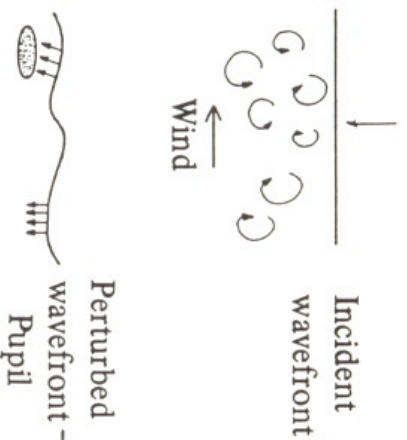


Image motion - the average slope of the wavefront at the pupil varies ("tip-tilt")

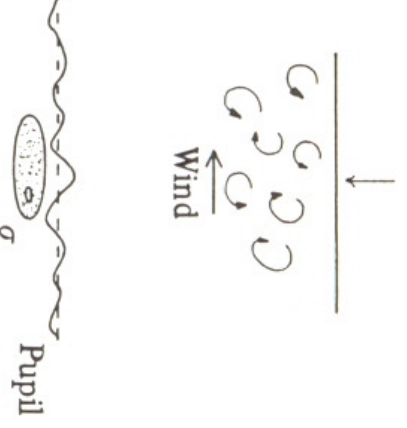
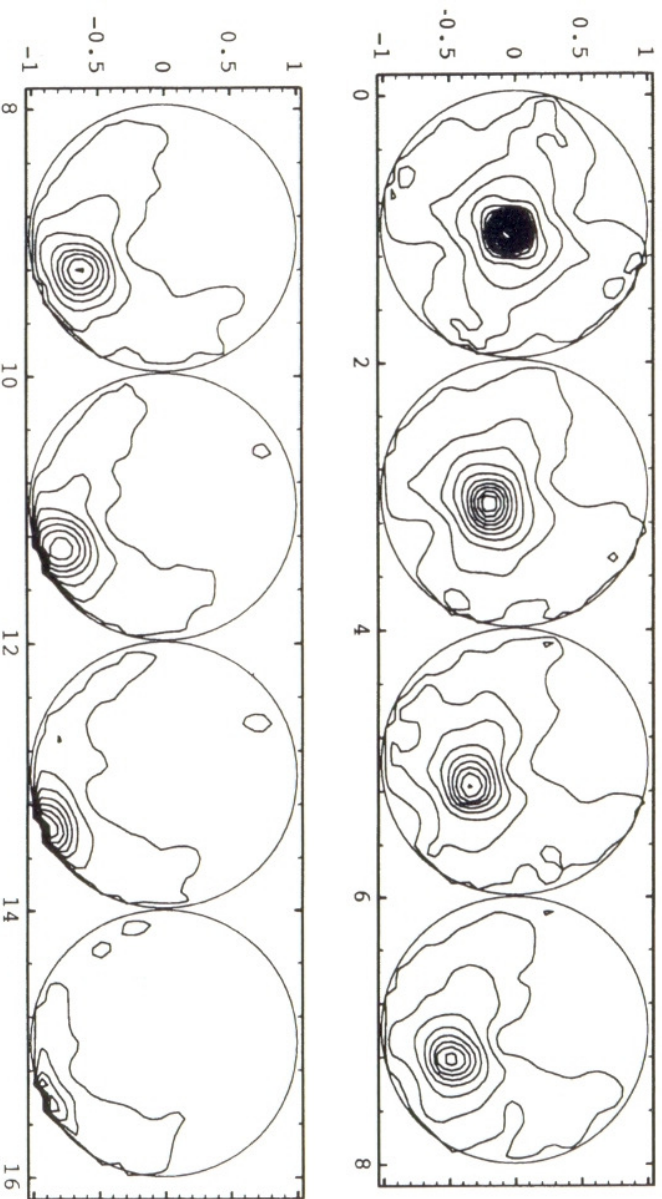


Image blurring - the spatial coherence of the wavefront is reduced ("seeing")

Turbulence Correlation Time τ_c

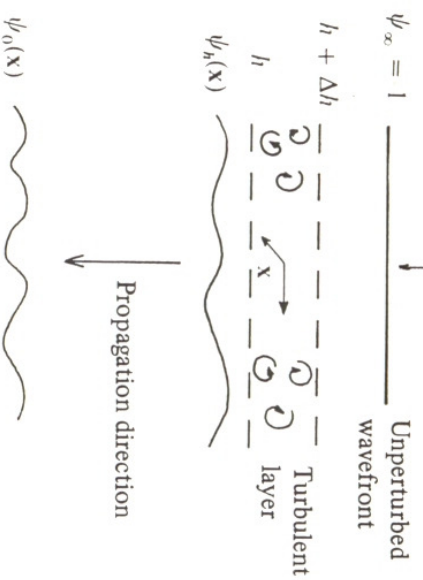
Often: time scales to generate turbulence \gg time for the turbulent medium to pass the telescope aperture (\leftarrow wind speed).
 \rightarrow correlation time τ_c .



Motion of a "frozen" patch of atmosphere across the 3.6m telescope aperture.
 Pictures by E. Gendron (1994)

Wavefront Perturbations

Consider a monochromatic, plane wave $\psi_\infty = 1$ which passes through a turbulent layer of thickness Δh .



Then the spatial correlation function of the wave across the telescope aperture is: $\langle \psi_h(x + \xi) \psi_h^*(\xi) \rangle = \exp(-1.45k^2 C_n^2 \Delta h x^{5/3})$, where $k = 2\pi/\lambda$ and the correlation length x_c is:

$$x_c \approx (1.45k^2 C_n^2 \Delta h)^{-3/5} = \left(1.45 \left(\frac{2\pi}{\lambda} \right)^2 C_n^2 \Delta h \right)^{-3/5} \propto \lambda^{6/5}$$

The Fried Parameter r_0

Related to the correlation length x_c is the so-called Fried parameter r_0 . It is the radius of the spatial coherence area:

$$r_0(\lambda) = 0.185 \lambda^{6/5} \left[\int_0^\infty C_n^2(z) dz \right]^{-3/5}$$

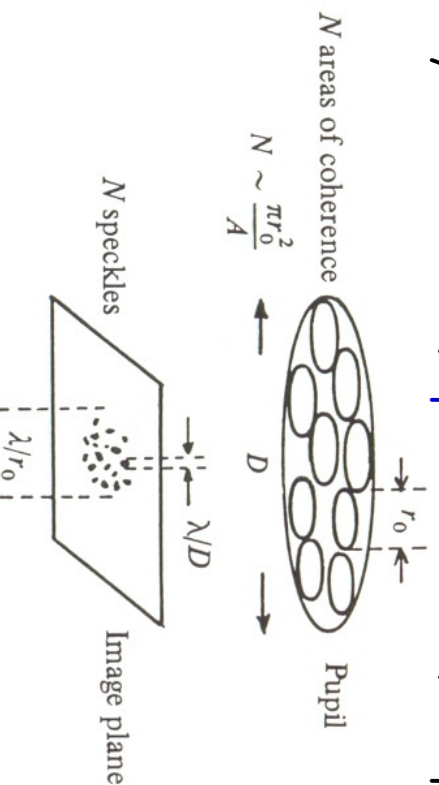
Note that r_0 increases as the 6/5 power of the wavelength and decreases as the -3/5 power of the air mass.

Another "definition" is that r_0 is the average turbulent scale over which the RMS optical phase distortion is 1 radian.

The angle $\Delta\theta = \frac{\lambda}{r_0}$ is often called the seeing.

Short Exposures through Turbulence

Random intensity distribution of speckles in the focal plane:



The observed image from some source is given by the convolution of I_0 with the MTF or pupil transfer function $T(\omega)$:

$$I(\theta) = I_0(\theta) * T(\theta) \quad \text{or} \quad \langle |I(\omega)|^2 \rangle = |I_0(\omega)|^2 \langle |T(\omega)|^2 \rangle$$

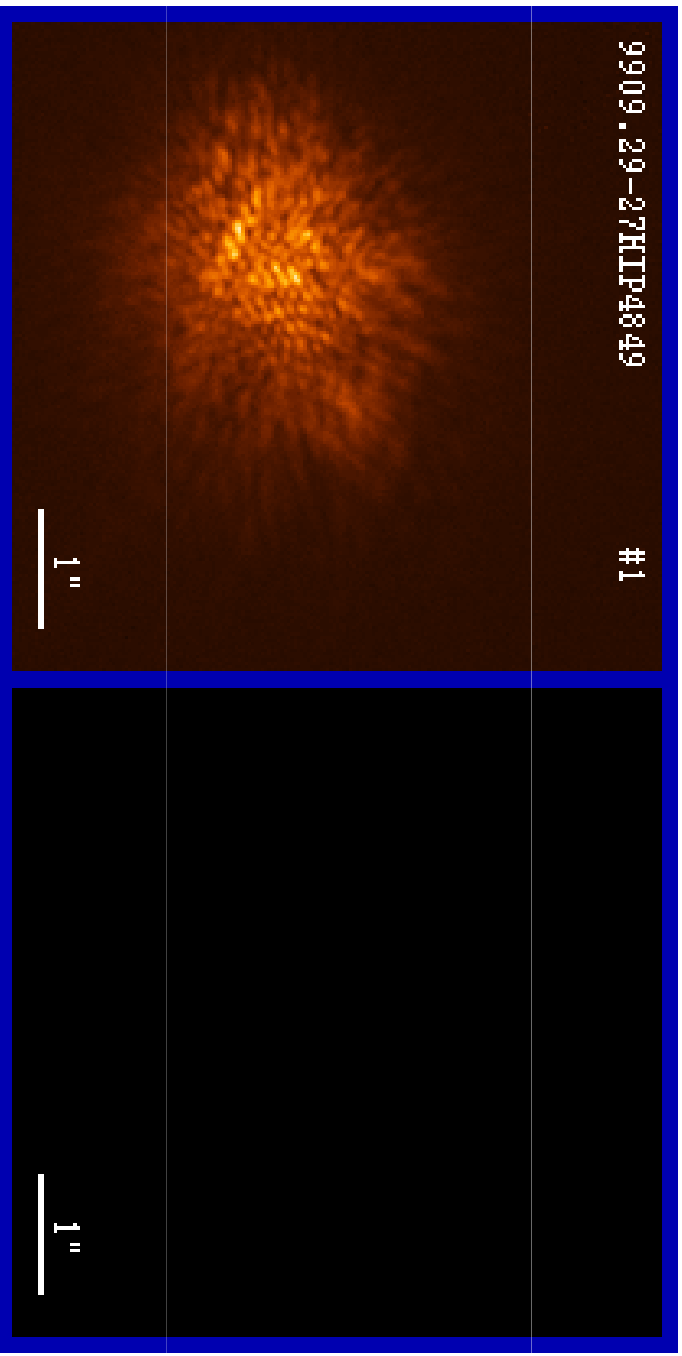
If a point source is observed as reference through the same r_0 we can

calculate: $|I_0(\omega)| = \left(\frac{\langle |I_0(\omega)|^2 \rangle_{obs}}{\langle |T(\omega)|^2 \rangle_{obs}} \right)^{1/2}$

This is called speckle interferometry.

Speckle Interferometry

Example: Real-time bispectrum speckle interferometry: 76 mas resolution.
<http://www.mpifr-bonn.mpg.de/div/ir-interferometry/movie/speckle/specklemovie.html>



Several related techniques do exist, e.g., Shift-and-add, Lucky Imaging, bispectrum analysis, Aperture masking, Triple correlation, ...

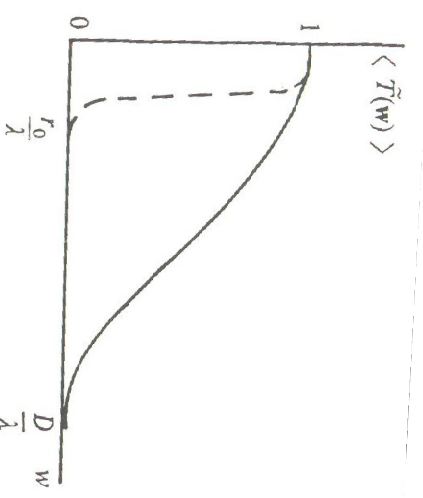
Long Exposures through Turbulence

When $t_{\text{int}} \gg \tau_c$ the image is the mean of the instantaneous intensity:

$$I(\theta) = \langle I_0(\theta) * T(\theta, t) \rangle$$

With the mean modulation transfer function (MTF):

$$\langle \tilde{T}(\omega) \rangle \approx \exp[-1.45k^2 C_n^2 \Delta h (\lambda \omega)^{5/3}]$$



→ The image is smeared or **spatially filtered** (loss of high spatial frequencies).

The angular dimension now has order of λ/r_0 rather than λ/D .

In other words:

As long as $D \gg r_0$, bigger telescopes will *not* provide sharper images.